Regional Setting, Representative Transects and Petroleum Potential of the Orphan Basin, offshore Newfoundland and Labrador

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Abstract

The Orphan Basin region, situated north of the Grand Banks of Newfoundland has undergone repeated extensional episodes. Crustal stretching of the mainly Paleozoic platform started in Late Triassic to Early Jurassic. Subsequent extensional episodes lasting until the Paleocene deepened the basin and enlarged it, by westward (landward) propagation of rifting. Seismic interpretation integrated with potential field data initially focused on two main regional transects and then extended to the entire basin, provides a regional understanding of the basin’s structural setting and evolution and indicates the presence of a viable petroleum system. Mapped hydrocarbon traps include large extensional anticlines which were also modified by compression. The presence of Jurassic source rock in the East Orphan Basin is convincingly supported by seismic correlation to source rock intervals within the adjacent Jeanne d’Arc and Flemish Pass basins, and by paleo-oceanic correlation with the Porcupine Basin off the western coast of Ireland.

Introduction

Newfoundland and Labrador’s divergent margin contains numerous Paleozoic and Mesozoic basins and sub-basins (Figure 1). The Orphan Basin is a wide continental rift located northeast and adjacent to the Jeanne D’Arc Basin from which approximately 400,000 barrels per day is currently being produced from three oil fields.

Figure 1. Distribution of sedimentary basin offshore Atlantic Canada (modified from C-NLOPB GSC, and Enachescu et al., 2005). Orphan Basin is shown with blue ellipse. Current exploration and production licences are shown in dark green.
With only seven wells and covering an area of 160,000 km², the Orphan Basin is very lightly explored and is currently the subject of an intense exploration effort by Chevron, ExxonMobil, Imperial Oil and Shell. The Basin is bounded by the Bonavista Fault zone to the west, Cumberland Belt Transfer Fault Zone (CBTZ) to the south, Continent-Ocean Boundary (COB) to the east and the Charlie-Gibbs Transfer Fault Zone (CBTZ) to the north (Figure 2).

**Figure 2.** Orphan Basin location map and basin boundaries. Bathymetry and 2003 ODP Sites 1276 and 1277, leg 210, after Tucholke, Sibuet et al., 2004. Tectonic framework from Enachescu et al., 2004a. Hibernia giant oil field and DSDP Site 111 are also shown. Annotations are CBTZ = Cumberland Belt Transform Fault Zone, CGTZ = Charlie Gibbs Transform Fault Zone, COB = interpreted Continent-Ocean boundary. The two main seismic transects researched and discussed in this report are also shown in Figures 3 and 4. Bathymetric highs are Grand Banks, Newfoundland’s margin shelfal part, Flemish Cap and Orphan Knoll (after Enachescu et al., 2005).

Water depth varies between 200m on the shelf, 2000-3000m within the eastern side of the basin and 4000m close to the COB. The Orphan Basin is a part of the complex network of interconnected basins and sub-basins located offshore Newfoundland and Labrador that evolved during the Mesozoic lithospheric extension, continental rifting, and continental breakup associated with the opening of the North Atlantic Ocean (Figure 1).

Various basins within this system are separated by basin-bounding faults, sediment-covered highs or by basement ridges devoid of sediments. An integrated geophysical interpretation was conducted over the Orphan Basin using modern, pre-stack time-migrated, 12-second record length, seismic data donated to Memorial University of Newfoundland by GSI of Calgary. The seismic interpretation was supplemented by public domain satellite gravity data and by the Geological Survey of Canada’s (GSC) suite of free-air gravity and total-field magnetic data base.
Seismic markers were correlated to 7 exploration wells located in the Orphan Basin and to 3 proximal wells within the Flemish Pass and Jeanne d’Arc basins. Two main seismic transects were used to illustrate the structural framework and the seismic-stratigraphic mega-sequences present in the basin.

**Geodynamic evolution**

Prior to the initiation of Mesozoic rifting the area of study was part of a broad Palaeozoic sedimentary platform within the Avalon terrane of the Appalachian orogenic system. Seismic stratigraphic relationships indicate that the first phase of rifting began in the Triassic, and affected only the eastern side of the Orphan Basin (“East Orphan Basin”) (Figures 3 to 5).

![West-East regional interpreted regional seismic section](image)

**Figure 3.** West-East regional interpreted regional seismic section (courtesy of GSI) across the Orphan Basin showing main structural and tectonic elements. Orphan Basin is a wide rift area formed by repeated extension (shown by red arrows) and transtension (shown by circles with direction of movement). Depth to water bottom is shown. Annotations are: Bsm = undifferentiated economic Basement, Tr = Triassic, J = Jurassic, K = Cretaceous; BT = Base Tertiary Unconformity.

This initial northeast-southwest, narrow rift expanded during Late Triassic-Early Jurassic within the Tethys rift system which extended from the Gulf of Mexico to the Barents Shelf and Northern Europe. After a long thermal-subsidence stage, the newly formed rift basin was reactivated, enlarged, and deepened during the Late Jurassic-Early Cretaceous Atlantic rifting phase. At that time, the East Orphan Basin was connected with other North Atlantic areas of known source rock and reservoir deposition, including: the oil-prolific Jeanne d’Arc Basin to the south, the source rock-proven Flemish Pass Basin to the southeast, and the petroliferous Porcupine Basin to the northeast.

Extension and minor transtension, continued during the Aptian-Albian (Labrador) rift phase and into several Late Cretaceous and Tertiary extensional episodes. As a result, the architecture of
the basin is dominated by alternating ridges of basement blocks overlain by sediments and deep sub-basins, which are predominantly oriented northeast-southwest or north-south, and reminiscent of the Basin and Range structural style seen in the south-western United States. The ridges within the Orphan Basin are very large and have comparable size to the Central Ridge forming the eastern boundary of the Jeanne d’Arc Basin. Several of the individual sub-basins within the Orphan Basin are of comparable size to Jeanne d’Arc Basin itself (Figures 3 to 5).

Figure 4. Interpreted NW-SE regional seismic section (courtesy of GSI) in East Orphan Basin. Annotations are: T = Tertiary, K = Cretaceous, J = Jurassic, Tr = Triassic interpreted seismic sequences. Pz = Paleozoic and Pc = Precambrian and Bsm = undifferentiated economic basement. Seismic markers are: Top of Economic Basement = yellow, Top Evaporites sequence (near Top Triassic) = blue, Top Jurassic = green and Top Cretaceous = pink. Mantle upwelling also interpreted from gravity data may represent the axis of the initial Late Triassic - Early Jurassic failed rift.

Figure 5. Time Structure of Economic Basement map (workstation generated) over the Orphan Basin included interpreted basin bounding Bonavista and West Sail faults. Red and yellow are high areas outside East Orphan Basin; within the basin green and light blues are local highs and dark blue, purple and dark purple are local lows; COH = Central Orphan High. Arrows indicate depression where seaway communication between the East Orphan and adjacent basins was possible during Triassic and Jurassic. White Sale Fault zone is made of many en-echelon segments.
 Structural Setting

Taking into account tectono-structural factors and area’s petroleum potential, the Orphan rifted area can be subdivided into (Figure 3 to 5):

1) An older East Orphan Basin situated in deep water (1,500-3,000m), which is a Tethys rift stage remnant with Jurassic (and probably Triassic) and Cretaceous sedimentary fill, and is likely to be oil and gas prone, and

2) A younger West Orphan Basin situated in shallower water (1000-1,500m), which evolved during the North Atlantic and Labrador rift stages containing mostly Cretaceous sedimentary fill, and which is likely to be gas prone.

The Tertiary cover is thick (4 km) over the W. Orphan Basin and relatively thin (2 km) over the E. Orphan Basin (Figure 6).

Figure 6. Tertiary Isopach map of the Orphan Basin and environs showing thick Tertiary cover in the West Orphan Basin, intermediate thickness cover in the East Orphan Basin, and thin Tertiary cover over the Central Orphan High and Flemish Pass/Cap area. Exploration licenses are also shown.

The two main rift basins, East and West Orphan, are separated by a major crustal fault zone, the White Sail fault which dips eastward and penetrate deeply into the upper crust. The N020° linear, tilted faults blocks identified within the West Orphan Basin run perpendicular to the flow-lines of the Flemish Cap motion during the M25-M0 period identified by magnetic
Paleo-reconstruction suggests that Flemish Cap behaved as a large, monolithic continental block that underwent a clockwise rotation of 43°, during the Late Triassic-Early Tertiary, which was accompanied by an apparent south-eastward displacement of more than 200 km. The associated stretching of the upper crust of the East Orphan Basin is observed to have a $\beta = 2.5$ (Figure 7). Regional seismic data suggest that trans-tensional movements, although hard to identify on reflection data, played an important role in the evolution of the basin and continued until the Paleocene. Beyond the Orphan Knoll-Flemish Cap “Outer Ridge” lineament lies a true
divergent margin basin showing little deformation within the sedimentary sequence and which overlies a relatively wide continent-ocean transition zone.

**Petroleum Potential**

Excluding the Orphan Knoll which has a thin Mesozoic section, and several of the elevated basement blocks devoid of Mesozoic sediments, the entire Orphan Basin has hydrocarbon potential. However, neither major shows, nor mature source rocks were encountered in the Orphan Basin in any of the seven wells in the exploration cycle that ran from 1975-1985. Nonetheless, good sandstone reservoirs were found in the Late Cretaceous and Tertiary formations. Additional evidence of reservoir potential was encountered in DSDP #111 on the Orphan Knoll which cored Bajocian-aged alluvial sandstones, the older syn-rift Mesozoic rocks drilled in the basin (Figure 2).

Late Jurassic paleogeographic reconstructions show the East Orphan rifted area as part of a larger regional rift system that included the basins of the Grand Banks and west Irish margin, where the presence of high quality Jurassic oil and gas source rocks has have been proven by drilling (Figure 7). Additionally, Middle to Late Cretaceous (post-Aptian) seaway extended from the Grand Banks into the Orphan Basin, and branched into northern Europe and the Labrador area (where deposition of Bjarni Formation and Markland shale gas-generating source rocks occurred). It is also noteworthy that Middle Cretaceous (Aptian - Turonian) source rocks account for more than 25% of world’s petroleum reserves and that the entire Orphan Basin contains significant intervals of Middle Cretaceous marine rocks.

![Figure 8](image)

**Figure 8.** Major exploration leads in the East Orphan Basin and environs shown on the Base Tertiary Depth Structure map and indicating active Exploration Licenses
Based on the age of basin fill, the presence of source rock, the timing of maturation, and seismic character correlation to adjacent basins, the Orphan Basin is predicted to be primarily oil-prone in its eastern part, and largely gas-prone in its western part. The East Orphan Basin (Figures 3 to 5, 8 and 9) is under active exploration, while there are no blocks currently under license in the West Orphan basin.

Figure 9. Seismic section (courtesy of GSI) showing Mesozoic basin continuity between the East Orphan Basin and Jeanne d’Arc Basin/Central Ridge area, where source rocks have been drilled in Panther P-52 well. The J = Jurassic interval can be correlated from southwest on the Central Ridge to northeast into the Orphan Basin.

Important points to consider regarding the Orphan Basin petroleum potential include the following:

1. Seismic data indicates that presently there is no total separation between the Orphan Basin and its southern neighbours - the Jeanne d’Arc and Flemish Pass basins (Figure 5). Several Jurassic-aged sediment-filled troughs show communication between basins across the Cumberland Belt fault zone. These connected areas have existed since the Jurassic, and possibly earlier. In its northern part, the Flemish Pass Basin appears as a higher terrace of the larger Orphan Basin.

2. Several key wells, Panther P-52, Baccalieu I-78, and Mizzen L-11, encountered thick Kimmeridgian-age source rocks (Egret Member of the Rankin Formation) They are located up-dip of and in areas connected to the Orphan Basin (e.g. Figure 9). Oil was also detected on logs by Petro-Canada et al. at Mizzen L-11. The Egret Member sequence penetrated in these wells can be easily correlated along seismic lines running into the licensed area of the Orphan Basin.

3. Structural closures with areas between 200 and 400 km² have been mapped in the East Orphan Basin. They are extensional anticlinal modified by transtension and
inversion. Not all of them lie within the area licensed for exploration and may thus become available in future landsales (Figures 3, 4, 8 and 10).

Figure 10. Seismic section (courtesy of GSI) showing a Late Jurassic-Early Cretaceous exploration lead on the western side of the East Orphan Basin. Annotations are: Bsm = undifferentiated economic Basement, Tr = Triassic, LJ = Late Jurassic, EK = Early Cretaceous,; LK = Late Cretaceous, T = Tertiary. Location on Base Tertiary structural map is shown in insert.

4. Seismic data show a series of complex features that cannot be explained strictly by dip-slip extension. Structural features vary from small inversions on fault planes to complex faulted anticlines and flower structures. They can be best described as compression modified extensional structures (CMES). In the absence of plate convergence in the area, mechanisms such as: oblique-extensional slip, transform faulting, microplate interaction, megaplate realignment, rotation of the extensional vector, post-rift gravity detachment and ridge push can be considered as sources of the inferred compressional stress.

5. As indicated by seismic character and amplitude changes, there is a high-probability of coarse siliciclastics being present in the basin. These are interpreted to be derived from the Bonavista Platform, Flemish Cap, Orphan Knoll, southern Labrador and Greenland areas, and the numerous local basement highs. Good-to-excellent Cretaceous sandstone reservoirs have been drilled in the Bonavista C-99, Linnet E-63, and Blue H-28 wells and in Flemish Pass and Jeanne d’Arc basins. Carbonate reservoirs of the Paleozoic “basement” have also shown reservoir properties in wells drilled in the Hopedale Basin, offshore Labrador.

6. Basin margin and basin-floor siliciclastic fans of Late Cretaceous and Tertiary age were derived from platforms and ridges and may be charged with hydrocarbons from postulated Albian, Late Cretaceous (equivalents of Bjarni Formation and Markland Shale of Labrador basins) or Tertiary source rocks. They are the most obvious undrilled play-types in the West Orphan Basin.

7. Numerous shallow and deep direct hydrocarbon indicators (including dramatic gas chimneys and at least one very definitive flat spot) are seen on modern 2D seismic
lines. One explanation for the presence of enigmatic deepwater seabed geomounds on seismic data identified on the Orphan Knoll (Enachescu, 2004) considers the possibility of a concentration of marine life near hydrocarbon seeps along deep-seated faults.

During the Late Jurassic - Early Cretaceous Atlantic riftting phase, the East Orphan Basin was connected with Jeanne d’Arc, Flemish Pass basins to the south and Porcupine and Rockall (eastern part) basins to the northeast. An excellent, oil-prone source rock was deposited in the neighbouring Jeanne d’Arc Basin that charged world-class hydrocarbon accumulations such as the Hibernia, Terra Nova, White Rose, and Hebron-Ben Nevis fields. Analyses of oils from the Jeanne d’Arc, Flemish Pass, Porcupine, and Celtic Sea basins indicate genetic geochemical links confirming geotectonic connection between conjugate margin basins.

Large 3D surveys were carried out in the past two years by an exploration partnership that includes Chevron Canada, ExxonMobil, Imperial Oil and Shell Canada that are focussed on eight large exploration licenses awarded in the 2003 land sale in the East Orphan Basin. Exploration drilling in what will be record water depths for Canada is expected to begin during 2006-2007.

Conclusion

A regional seismic study anchored on several plate margin-wide transects was carried out by the Basin Analysis group at Memorial University of Newfoundland, using PRAC, NSERC and Industry funds. Based on tectonic reconstruction considerations, it is proposed that the Orphan Basin can be subdivided into two large sub-basins which are divided along the White Sail Fault Zone. In consideration of seismic correlations with adjacent basins it would appear that the East Orphan Basin is likely to contain Kimmeridgian oil and gas prone source rock as has been encountered in the Jeanne d’Arc and Flemish Pass Basins. The West Orphan Basin is more likely to contain Cretaceous gas-prone source rocks, the same as the ones that have been encountered on the Labrador Shelf. It is also postulated that during basin evolution, the Flemish Cap continental block underwent clockwise rotated and was displaced south-eastward approximately 200km.

Large basement-cored ridges and domes overlain by syn-rift sediments or rollover anticlines in the downthrown blocks of major faults, some modified by compression, provide very large structural traps within the East Orphan Basin, which are likely to be sourced Kimmeridgian shales (Egret Member equivalent). Platform and ridge-derived siliciclastic fans of Late Cretaceous and early Tertiary age, with hydrocarbons sourced from a postulated Albian or Late Cretaceous-early Tertiary source rocks present a very significant undrilled play in the West Orphan Basin. Petroleum potential of the East Orphan Basin will be tested by a deepwater drilling program (3-8 wells) that is expected to begin in 2006-2007.

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Further Reading


Thompson, T., 2003, Preliminary findings on basin architecture, segmentation and inversion on a passive margin offshore Newfoundland: Expanded Abstract, Presented at Calgary CSPG/CSEG Joint Conference.


**Author Biography**

**Michael Enachescu** is Husky Energy Senior Fellow in Exploration Geophysics at Memorial University of Newfoundland, an Associate Professor at the Department of Earth Sciences, Pan-Atlantic Petroleum Systems Consortium (PPSC) and Oil and Gas Development Partnership (OGDP) and an advisor to Palo Alto Investors group (PAI) and to several oil companies, seismic contractors and scientific panels. He worked in resource exploration and geophysical research in Europe and after 1981 as a petroleum explorationist in Calgary. He has been involved with major exploration drilling programs in the Grand Banks, Scotian Shelf and Slope, Labrador Sea, Arctic, Beaufort Sea first with Suncor Resources, Trillium/Mosbacher and from 1984 to 2003 with Husky Energy. Michael was a member of the regional mapping, discovery and delineation teams and a contributor to the Development Plan Applications for Terra Nova and White Rose fields, offshore Newfoundland. Michael was a member of the Scientific Committee of LITHOPROBE, a member of the Site Survey Panel of ODP and IODP (1994-2003) and a volunteer with CSEG (2nd VP, Technical Chair GeoCanada 2000 convention and repeatedly annual meeting session chair) and with many other professional societies and charity organizations. Since at Memorial (fall 2003), Michael is teaching Atlantic Geology, Rift Tectonics, Marine Seismic and Seismic Interpretation courses and conducts research with a group of 12 graduate students in structure and tectonics and petroleum systems of Newfoundland and Labrador’s offshore basins. Michael has extensively published on the structural setting and petroleum geology of Atlantic Canada and received the 1999 CSEG Meritorious Award and is the 2006 CSEG Distinguished Lecturer. Michael is a member of CSEG, SEG, EAEG, CGU, AAPG, CSPG, RGS, and a P. Geoph. with APEGGA.